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### **Study Of Weather Analysis Via Geometric Modeling**

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#### **ABSTRACT**

Weather phenomena such as sunbeams, clouds of all sorts, precipitation (ranging from light sprinkle to rain to hail and snow), fog, lightning, wind, humidity, and hot and cold conditions can all be part of much-larger-scale weather systems. The weather systems are cyclones (low pressure systems) and anticyclones (high-pressure systems) and the associated warm and cold fronts. The Weather condition of the atmosphere at a given time and place, usually expressed in terms of pressure, temperature, humidity, wind, etc. Also, the various phenomena in the atmosphere occurring from minutes to months. In this paper, we have defined and studied weather map analysis for numerical weather prediction via geometric modeling.

**KEYWORDS:** Geometric modeling, Weather phenomena, atmosphere, cyclones and anticyclones.

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## **1. INTRODUCTION**

An important concept to hold is how weather patterns and the kinds of weather that occur relate to climate. We submit to this connection as the weather machine because of the way the weather helps drive the climate system. It is the sum of many weather phenomena that determines how the large-scale general circulation of the atmosphere. That is, the average three-dimensional structure (Geometric modeling) of atmospheric motion. Actually works and it is the circulation that essentially defines climate. This intimate link between weather and climate provides a basis for understanding how weather events may change as the climate changes.

We generally think of the fluctuations in the atmosphere from hour to hour or day to day as weather. Weather is described by such elements as temperature, air pressure, humidity, cloudiness, precipitation of various kinds, and winds. Weather occurs as a wide variety of phenomena ranging from small cumulus clouds to giant thunderstorms, from clear skies to extensive cloud decks, from gentle breezes to gales, from small wind gusts to tornadoes, from frost to heat waves, and from snowflurries to torrential rain. Many such phenomena occur as part of much larger-scale organized weather systems which consist, in middle latitudes of cyclones (low pressure areas or systems) and anticyclones (high pressure systems) and their associated warm and cold fronts. Tropical storms are organized, large-scale systems of intense low pressure that occur in low latitudes. If sufficiently intense these become hurricanes, which are also known as typhoons or tropical cyclones in other parts of the world.

## **2. WEATHER ANALYSIS**

The glossary defines “analysis” as a comprehensive examination of something. It is a process that breaks the object of the analysis into parts so that each part can be examined in detail. This process determines its nature, function, and other characteristics. In operational weather analysis the object is the atmosphere, particularly, the troposphere, where day-to-day weather occurs. The parts that we need to examine are the various weather parameters that are naturally used to describe the atmosphere, temperature, humidity, wind speed and direction, atmospheric pressure, clouds, precipitation, etc.

We generally do our analysis at one specific time in order to get a glance of what is happening in the atmosphere at that moment. However, analysis over time is often useful to provide a better picture of how things are changing. In weather analysis and forecasting, I examined a wide variety of weather data throughout the troposphere by integrating several observing systems during the weather analysis process. We have combine all this information into a three-dimensional image (geometric

modeling) of what is producing the current weather by correlating what is observed with meteorological background.

Following tools are used in weather map analysis (*Source: Surface weather analysis- Wikipedia.*)

## **Pressure Centers**

Centers of surface (high-pressure and low-pressure) areas that are found within closed isobars on a surface weather analysis are the absolute maxima and minima in the pressure field, and can tell a user in a glance what the general weather is in their vicinity.

### **Low Pressure**

Low-pressure systems, also known as cyclones, are located in minima in the pressure field. Rotation is inward at the surface and counterclockwise in the northern hemisphere as opposed to inward and clockwise in the southern hemisphere due to the Coriolis force.

### **High Pressure**

High-pressure systems, also known as anticyclones, rotate outward at the surface and clockwise in the northern hemisphere as opposed to outward and counterclockwise in the southern hemisphere.

## **Fronts**

Fronts in meteorology are the leading edges of air masses that have a density, air temperature, and humidity different from the air mass it is invading. When an air mass passes over an area, it is marked by changes in temperature, moisture, wind speed and direction, atmospheric pressure, and often a change in the precipitation pattern. The change is called a front, whether it is warm or cold.

### **Cold Front**

A cold front's location is at the leading edge of the temperature drop-off, which in an isotherm analysis shows up as the leading edge of the isotherm gradient, and it normally lies within a sharp surface trough. Cold fronts can move up to twice as fast as warm fronts and produce sharper changes in weather, since cold air is denser than warm air and rapidly lifts the warm air as the cold air moves in.

### **Warm Front**

Warm fronts mark the position on the Earth's surface where a relatively warm body of air has displaced colder air. The temperature increase is located on the equatorward edge of the gradient in isotherms, and lies within broader low pressure troughs than is the case with cold fronts. Warm fronts move more slowly than do the cold fronts because cold air is denser, and harder to displace from the Earth's surface.

## **Occluded Front**

An occluded front is formed during the process of cyclogenesis when a cold front overtakes a warm front. The cold and warm fronts curve naturally poleward into the point of occlusion, which is also known as the triple point in meteorology.

## **Stationary Fronts and Shear lines**

A stationary front is a non-moving boundary between two different air masses, neither of which is strong enough to replace the other. They tend to remain in the same area for long periods of time, usually moving in waves. There is normally a broad temperature gradient behind the boundary with more widely spaced isotherms. A wide variety of weather can be found along a stationary front, but usually clouds and prolonged precipitation are found there. Stationary fronts will either dissipate after several days or devolve into shear lines, but can change into a cold or warm front if conditions aloft change causing a driving of one air mass or the other.

## **Dry Line**

The dry line is the boundary between dry and moist air masses east of mountain ranges with similar orientation to the Rockies, depicted at the leading edge of the dew point, or moisture, gradient. Near the surface, warm moist air that is denser than warmer, drier air wedges under the drier air in a manner similar to that of a cold front wedging under warmer air. When the warm moist air wedged under the drier mass heats up, it becomes less dense and rises and sometimes forms thunder storms.

## **3. NUMERICAL WEATHER PREDICTION MODEL**

Since the early 20th century numerical weather prediction (NWP) has increasingly become one of the most important and complicated problems of modern science. With the advent of computers, increased observations, and progress in theoretical understanding, numerical models were developed. Since then, such models are playing an increasing role in understanding and predicting weather and climate and have been a driving force in the advancement of the meteorological sciences. Numerical models are a mathematical representation of the earth's climate system including the atmosphere, ocean, cryosphere and land, among others. The models divide the area of interest into a set of grids and then make use of observations of variables such as surface pressure, winds, temperature and humidity at numerous locations throughout the globe. The observed values are then assimilated and used by the model to predict future evolution of the earth's weather and climate. In the mid 20th century, models evolved from a simple model with a single atmospheric layer to a multi-layer primitive equation model capable of predicting cyclone development (Stensrud, 2007). Numerical Weather Prediction (NWP) is the science of predicting the weather using

models of the atmosphere and computational techniques. Current weather conditions are used at the input of the mathematical models of the atmosphere to predict the weather.

Weather and climate models are based on the Navier-Stokes equations and moist thermodynamics, but substantial insights into the processes and mechanisms behind ubiquitous features of weather were obtained long before computing power reached the levels we take for granted today. These insights were obtained by applying systematic approximations to the Navier-Stokes equations, while retaining important properties such as conservation laws for energy and potential vorticity. Many decades of careful research have resulted in a deep understanding of what lies behind the choreography of weather, and this understanding is utilized in the design of numerical models. By incorporating this knowledge via elegant numeric's, we focus on the features that influence the weather for the week ahead, rather than expending computer power on all the eddying local gusts of wind that play little or no part in the evolution of the big picture.

We shall use the (Geometric Modeling) one-dimensional water equations as a vehicle to describe the salient mathematical ideas that help meteorologists to quantify some of the large-scale, often highly predictable, features of the atmospheric circulation. For brevity, we ignore thermodynamics in the description that follows, but the model is readily augmented to include the physics of heat and moisture (which are necessary if considering phenomena such as weather fronts), while retaining the key mathematical structures.

The one-dimensional equations (Geometric Modeling) on a domain of  $R^2$ , with local Cartesian coordinates  $(x, y)$ , and rotating with constant angular frequency  $f/2$ , are

$$\frac{Du}{Dt} - fv + g \frac{\partial h}{\partial x} = 0, \quad \frac{Dv}{Dt} + fu + g \frac{\partial h}{\partial y} = 0, \quad (3.1)$$

$$\frac{Dh}{Dt} + h \nabla \cdot \mathbf{u} = 0, \quad (3.2)$$

Where  $t$  denotes time,  $g$  is the acceleration due to gravity (a constant) and  $h(x, y, t)$  is the depth of the fluid. The horizontal velocity has two components  $\mathbf{u}(x, y, t) = (u, v)$ , and the Lagrangian, or material,

$$\text{derivative is } \frac{D}{Dt} = \frac{\partial}{\partial t} + \mathbf{u} \cdot \nabla. \quad (3.3)$$

The Lagrangian derivative represents the change of a dependent variable as we "follow the flow"; i.e. it is a directional along the trajectory of a fluid particle. The derivative  $\mathbf{u} \cdot \nabla$  is often referred to as the advection, or transport, term. If any variable  $A(x, t)$  satisfies  $DA/Dt = 0$ , then we say that  $A(x, t)$  is conserved in the Lagrangian sense. Such conservation laws are of fundamental importance in metrology and oceanography. The so-called semi-geostrophic equations amount to following approximation to the shallow water equations

$$\frac{Du_g}{Dt} - fv + g \frac{\partial h}{\partial x} = 0, \quad \frac{Dv_g}{Dt} + fu = g \frac{\partial h}{\partial y} = 0, \quad (3.4)$$

$$\frac{Dh}{Dt} + h\nabla \cdot u = 0, (3.5)$$

Where  $u_g$  and  $v_g$  are two components of the geostrophic wind  $g\nabla h = (fv_g, -fu_g)$ .

The geostrophic flow is parallel to constant h (in the context of the Navier –Stokes equations on a rotating domain, the geostrophic flow is parallel to the isobars of constant pressure). The difference between the shallow water equations and the semi-geostrophic equations is the replacement of the fluid velocity (u, v) with the geostrophic wind ( $u_g, v_g$ ) in the Lagrangian derivatives of u and v, while leaving the derivatives operator (3.3) and the continuity equation (3.5) unchanged. This is known as the geostrophic momentum approximation (Hoskins 1975), and it applies to flows in which the rate of change of momentum is much smaller than the Coriolis force. The advecting velocity (that is, the velocity appearing in the Lagrangian derivative  $u \cdot \nabla$ ) is not approximated in semi-geostrophic flows, such as jet streams and fronts, have two distinct length scale (for example, a weather front is a relatively sharp discontinuity between air masses, but fronts extend to distances of the order of 1000km along the interface between air masses). In essence, the geostrophic momentum approximation tells us that it is important to represent the advecting velocity accurately; while the quantity is being advected (e.g. the geostrophic wind) can be approximated.

Equations (3.4) and (3.5) conserve energy and the following form of potential vorticity

$$q = \frac{1}{h} \left( f + \frac{\partial v_g}{\partial x} - \frac{\partial u_g}{\partial y} + \frac{1}{f} \frac{\partial(u_g, v_g)}{\partial(x,y)} \right). (3.6)$$

In atmospheric dynamics (with thermodynamics included), potential vorticity is proportional to the vector dot product of vorticity and stratification that, following the flow, can only be changed by diabatic or frictional processes. Potential vorticity is a fundamental concept for understanding the generation of vorticity in cyclogenesis (the birth and development of a cyclone), especially along the polar front, and in analyzing flow in the ocean. The use of potential vorticity played a key role as a diagnostic in understanding the evolution of Hurricane Sandy.

The semi-geostrophic equations have played a major role in understanding the formation of fronts (front genesis), and the properties of the equations that facilitate such studies are revealed through a transformation of coordinates. Defining new coordinates (sometimes called geostrophic momentum coordinates)

$$\mathbf{X} \equiv (X, Y) \equiv (x + v_g, y - u_g) (3.7)$$

We find that (4) may be replaced by

$$\frac{DX}{Dt} = u_g \equiv (u_g, v_g). (3.8)$$

Hence the motion in these transformed coordinates is exactly geostrophic. The Jacobian is proportional to the potential vorticity (3.6)

$$\frac{\partial(X,Y)}{\partial(x,y)} = \frac{hq}{f} \quad (3.9)$$

The vector X may be expressed as the gradient of a scalar function P(x),

$$X = \left(\frac{\partial P}{\partial x}, \frac{\partial P}{\partial Y}\right), \quad (3.10)$$

Which, to within an arbitrary additive constant, is uniquely defined by

$$P(x, t) = \frac{1}{2}(x^2 + y^2) + \frac{gh(x,t)}{f^2}. \quad (3.11)$$

We note, using  $g\nabla h = (fv_g, -fu_g)$  together with (3.7) and (3.11), that (3.9) has the form of a Monge-Ampere equation for P, given q(x, y, t) and suitable boundary conditions

$$q = \frac{f}{h}(P_{xx}P_{yy} - P_{xy}^2). \quad (3.12)$$

When the Jacobian (9) is non-singular, we express x(X) by introducing a scalar function

$$R(X):x = \left(\frac{\partial R}{\partial X}, \frac{\partial R}{\partial Y}\right), \quad (3.13)$$

Where R is given to within an additive constant by

$$R(X) = x.X - P(x) \quad (3.14)$$

(Time is a parameter in this transformation). Equation (3.14) is the expression for the Legendre transformation between R(X) and p(X). Local singularities of this map can be interpreted as atmospheric fronts (Chynoweth and Sewell 1989).

The semi-geostrophic equations can be integrated in time using the conservation of potential vorticity, expressed in geostrophic momentum coordinates. We write the reciprocal of the potential vorticity,  $q^{-1}$  as

$$q(x, t)^{-1} \equiv \rho(X, t) = \frac{h(x,y)}{f}(R_{XX}R_{YY} - R_{XY}^2) \quad (3.15)$$

This may be expressed solely in terms of the geostrophic momentum coordinates by defining

$$\phi(x, t) = gh(x, t)/f^2 \text{ and } \Phi(X, t) = \frac{1}{2}(X^2 + Y^2) - R(X, t), \text{ then we note}$$

$$\frac{\partial \Phi}{\partial X} = \frac{\partial \phi}{\partial x} = X - x, \frac{\partial \Phi}{\partial Y} = \frac{\partial \phi}{\partial y} = Y - y.$$

Hence (3.15) may be written

$$\rho(X, t) = \frac{f}{g}(\Phi - \frac{1}{2}(\Phi_X^2 + \Phi_Y^2)) | \text{Hess}(\frac{1}{2}(X^2 + Y^2) - \Phi) |,$$

Where Hess (·) is the hessian matrix of the second derivatives of (·) with respect to X,Y.

We can show that  $\rho$  satisfies

$$\frac{\partial \rho}{\partial t} + \dot{X} \frac{\partial \rho}{\partial X} + \frac{\partial \rho}{\partial Y} = 0 \quad (3.16)$$

Where  $\dot{X} = f\left(\frac{\partial R}{\partial Y} - Y\right)$ .  $\dot{Y} = -f\left(\frac{\partial R}{\partial X} - X\right)$ . (3.17)

The integration begins by solving the dual Monge-Ampere equation (3.15) for R, and then using R in (3.17) can be expressed in Hamiltonian form (McIntyre and Roulstone 2002):

$$\dot{X} = -f \frac{\partial \Phi}{\partial Y}, \dot{Y} = f \frac{\partial \Phi}{\partial X}.$$

The semi-geostrophic equations involve two important types of geometry, symplectic geometry associated with the Hamiltonian structures and contact geometry associated with the Legendre transformation. These two geometries also play a fundamental role in the theory of the Monge-Ampere equation (Kushner et al. 2007) and, in turn, relate to complex geometries, such as Kaehler geometry. The transformation properties of the semi-geostrophic equations can be understood within the context of hyper-Kahlar geometry, and this geometry has been exploited in the study of a much broader class of models of cyclones (McIntyre and Roulstone 2002, Delahaies and Roulstone 2010).

The emergence of a complex structure, which at first sight is somewhat surprising given the classical nature of the fluid mechanics, occurs when the Monge-Ampere equation (3.12) is elliptic. This equation is elliptic when  $q > 0$ . the ellipticity is related to convexity condition on the energy and to the stability of the flow (Cullen 2006). The total energy, which is conserved by the shallow water equations (3.1), (3.2), is

$$E = \int (\frac{1}{2}(u^2 + v^2) + \frac{1}{2}gh) h dx dy. \quad (3.18)$$

This is a function of u, v and h, and the conditions for E to be minimized corresponds to the stability of a geostrophic flow, viewed as a solution of the unapproximated equations, to perturbations of the form

$$\delta u = f \delta y, \delta v = -f \delta x, \nabla \cdot (\delta x, \delta y) = 0. \quad (3.19)$$

The second variation of E is greater than zeros when the Hessian matrix of the Legendre function P is positive definite. This corresponds to the ellipticity of the Monge-Ampere equation (3.2). Introducing momentum coordinates,  $(X', Y')$ , via  $(u, v) \equiv f(y - Y', X' - x)$ , we can show that the perturbations (3.19) imply  $\delta \sigma = 0$ , where

$$\sigma = h \frac{\partial(x, y)}{\partial(X', Y')}. \quad (3.20)$$

If we define s distance  $d(x, X')$  between x and  $X'$  such that

$$d^2 = f^2((X' - x)^2 + (Y' - y)^2),$$

Then the energy functional can be rewritten

$$E = \int (\frac{1}{2}d^2(x, X') + \frac{1}{2}gh) h dx dy. \quad (3.21)$$



The proof of E can be uniquely minimized is established by showing that, given  $\sigma$  as a non-negative function of the momentum coordinates, there is a unique mapping from  $(X', Y')$  to  $(x, y)$  that minimizes E and satisfies (3.20). This is the starting point for expressing the energy minimization as a Monge optimal mass transport problem (Benamou and Brenier 1998; Cullen 2006; Villani 2008). Utilizing theorems on the regularity of solutions of the Monge-Ampere equation, together with optimal mass transport theory, it has been shown that the semi-geostrophic equations can be integrated for large times from suitable initial data. (Cullen and Roulstone, 1993) showed that a semi-geostrophic finite element numerical simulation of an Eady wave- the basic building block of cyclogenesis could be performed through multiple life-cycles, and the predictability of the key features of the system were remarkably robust to perturbations of the initial data.

#### **4. CONCLUSION**

Since the first forecast, the complexity and sophistication of numerical weather prediction Models have increased tremendously. The continued improvement in data assimilation and numerical models, and the continued availability of ever larger and faster computers have allowed numerical weather prediction to become more accurate. Most of the improvements in numerical models that occurred over the past 50 years can be categorized as improved numerical techniques, improved model resolution, or improved model physics. Development of different types of grids, grid-staggering and grid-spacing are intimately related to the improvement in the weather and climate prediction.

The target of numerical weather simulations is to calculate the state of atmosphere depending on time. That means the simulation has the purpose to calculate the velocity, density, pressure, temperature and humidity of every single point in the air. As it is not possible to regard every single point because of observational and computational limitations a two-dimensional or a three-dimensional manifolds is used for the ability of the NWP model to accurately represent atmospheric phenomena is based on three conditions; scientific knowledge, the availability of observational data, and computer processing abilities. If enough observational data is available and enough scientific knowledge is present then the limiting factor of an accurate forecast is the power of the processing computer.

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